International Workshop on Stochastic Weather Generators

Reduced flow models from a stochastic Navier-Stokes theorem

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PLAN

INTRODUCTION

STOCHASTIC FLUID DYNAMIC MODEL

REDUCED MODEL

RESULTS

CONCLUSION

Plan

INTRODUCTION

Why using stochastic fluid dynamic? Why a reduced model?

STOCHASTIC FLUID DYNAMIC MODEL

REDUCED MODEL

RESULTS

CONCLUSION

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INTRODUCTION

REDUCED MODEL 0000000

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WHY USING STOCHASTIC FLUID DYNAMIC?



- Usual advantages of stochastic methods
- Fluid dynamics, and especially geophysical fluid dynamics, is too complex to be solved at all scales
- Chaotic system
- Non-unique solution

WHY A REDUCED MODEL?



- Simplified simulations
- Decomposition of the physic



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INTRODUCTION

STOCHASTIC FLUID DYNAMIC MODEL

Classical Large Eddy approach Stochastic Reynolds-transport theorem Mass continuity Stochastic Navier-Stokes model

REDUCED MODEL

RESULTS

CONCLUSION

$$\frac{\partial \overline{T}}{\partial t} + \nabla \cdot \left(\overline{T} \overline{u} \right) = \nabla \cdot \left(- \overline{u' T'} \right)$$

$$\frac{\partial \overline{T}}{\partial t} + \nabla \cdot \left(\overline{T}\overline{u}\right) = \nabla \cdot \left(-\overline{u'T'}\right) = \frac{\sigma^2}{2} \triangle \overline{T}$$

$$\frac{\partial \overline{T}}{\partial t} + \nabla \cdot (\overline{T}\overline{u}) = \nabla \cdot (-\overline{u'T'}) = \frac{\sigma^2}{2} \triangle \overline{T} \rightarrow \text{Eddy diffusivity}$$

$$\frac{\partial \overline{T}}{\partial t} + \nabla \cdot (\overline{T}\overline{u}) = \nabla \cdot (-\overline{u'T'}) = \frac{\sigma^2}{2} \triangle \overline{T} \rightarrow \text{Eddy diffusivity}$$



STOCHASTIC REPRESENTATION OF THE REYNOLDS-TRANSPORT THEOREM

Notations :

- $\bullet \ dX_t = w(X_t, t)dt + \sigma(X_t, t)dB_t$
- *B_t* an I_d-cylindrical Wiener process (Da Prato & Zabczyk (1992))
- σ(., t) an Hilbert-Schmidt operator on (L²(ℝ^d))^d
 defined by its kernel ŏ(.,.,t)

Thus,

$$\sigma(x,t)dB_t \sim \mathbf{GP}(0,\delta(t-t')\sigma(x,t)\sigma^t(x',t)dt)$$

► Local covariance matrix $a(x,t) \stackrel{\triangle}{=} \sigma(x,t)\sigma^t(x,t) = \int_{\Omega} \breve{\sigma}(x,z,t)\breve{\sigma}^t(x,z,t)dz$

STOCHASTIC REPRESENTATION OF THE REYNOLDS-TRANSPORT THEOREM

Theorem V(t) being a material element,

$$d_t \int_{V(t)} q dx = \int_{V(t)} \left(d_t q + \nabla \cdot \left(q dX_t + q \sigma (\nabla \cdot \sigma)^t dt - \frac{1}{2} \nabla \cdot (aq)^t dt \right) \right) dx$$

Proof : Mémin (2014)

MASS CONTINUITY

$$d_t \int_{V(t)} \rho dx = d_t m(t) \\ = 0$$

Thus,

Theorem

$$0 = d_t \rho + \nabla \cdot \left(\rho dX_t + \rho \sigma (\nabla \cdot \sigma)^t dt - \frac{1}{2} \nabla \cdot (a\rho)^t dt \right)$$

Application : Mean field

$$\frac{\partial \mathbb{E}(\rho)}{\partial t} + \nabla \cdot \left(\mathbb{E}(\rho)w^*\right) = \nabla \cdot \left(\frac{1}{2}a\nabla \mathbb{E}(\rho)\right)$$

with

$$w^* = w + \sigma (\nabla \cdot \sigma)^t - \frac{1}{2} (\nabla \cdot a)^t$$

STOCHASTIC NAVIER-STOKES MODEL Theorem

If w has finite variations and $\int_0^t pdt'$ is replaced by $\int_0^t p'dt' + d\hat{p}$ then

$$\rho\left(\frac{\partial w}{\partial t} + (w \cdot \nabla)w\right) = \boldsymbol{\tau}(w) + \rho g - \nabla p' + f_V(w)$$
$$\rho(\sigma dBt \cdot \nabla)w = -\nabla d\hat{p} + f_V(\sigma)dBt$$

where

$$\begin{cases} \forall k, \tau_k(w) = \frac{1}{2} \left(\nabla \cdot (\nabla \cdot (\rho a w_k))^t - \nabla \cdot (\nabla \cdot (\rho a))^t w_k \right) \\ & -\rho \left((\nabla \cdot \sigma) \sigma^t \nabla \right) w_k \\ f_V(g) = \mu \left(\nabla^2 g + \frac{1}{3} \nabla (\nabla \cdot g) \right) \end{cases}$$

Proof : Mémin (2014)

STOCHASTIC NAVIER-STOKES MODEL

Applications :

- Large eddies simulation
- Uncertainty quantification
- Filtering

...

Mixing diagnostics

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INTRODUCTION

STOCHASTIC FLUID DYNAMIC MODEL

REDUCED MODEL

Principle Proper Orthogonal Decomposition Classical approaches Our approach Estimation of *a*

Results

CONCLUSION

PRINCIPLE

- Using both a model and data (*N* snapshots).
- Galerkin projection :

$$u(.,t) \in \mathbf{Span}(\phi_0,...,\phi_n)$$

▶ Coefficients of the decomposition, b_i, are time-dependent.
 ▶ ∂u/∂t = I + L(u) + C(u, u) (a PDE) becomes :

$$\forall i, \frac{db_i}{dt} = \left(\int_{\Omega} \phi_i \cdot I\right) + \sum_{p=0}^n \left(\int_{\Omega} \phi_i \cdot L(\phi_p)\right) b_p + \sum_{p,q=0}^n \left(\int_{\Omega} \phi_i \cdot C(\phi_p, \phi_q)\right) b_p b_q \qquad (\text{ODEs})$$

PROPER ORTHOGONAL DECOMPOSITION

- Analogue to PCA apply to (u(., t₁), ..., u(., t_N)): Analogue of variables : Values on points of the space Analogue of realizations : Snapshots
- ► Modes $\phi_i(x)$ are sorted regarding an energetic criterion

$$u'(.,t) \stackrel{\triangle}{=} u(.,t) - \underbrace{\bar{u}}_{\triangleq \phi_0} \approx \sum_{i=1}^N b_i(t)\phi_i \approx \sum_{i=1}^n b_i(t)\phi_i$$

PROPER ORTHOGONAL DECOMPOSITION







CLASSICAL APPROACHES

$$\forall i, \frac{db_i}{dt} = i_i + l_{\cdot i}b_{1:n} + b_{1:n}^t c_{\cdot \cdot i}b_{1:n}$$

Problem :

Non-linear systems require too many modes.

Classical solutions : Eddy viscosity

$$l._i \leftarrow \frac{\nu}{\nu_m} l._i$$

- ▶ Constant *ν* (Aubry *et al.* (1988))
- Modal ν_i (Rempfer & Fasel (1994))
- ν or $\nu_i \propto ||b_{1:n}||_2$ (Östh *et al.* (2014))
- $\nu \propto f(\|b_{1:n}\|_2^2)$ (Protas *et al.* (2014))

OUR APPROACH

Our approach :

 dX_t = wdt + σdB_t
 w = ∑_{i=0}ⁿ b_iφ_i (in the truncated subspace)
 ∑_{i=n+1}^N b_iφ_idt a realization of σdB_t (in the complementary "small-scale" subspace)

ESTIMATION OF *a*

We look for a decomposition of $a = \sum_{i=0}^{n} b_i(t) z_i(x)$.

- Reduced subspace with coherent time scales
- No need to reconstruct a
- Autonomous system

Indeed, $\frac{\partial w}{\partial t} = I + L(w) + C(w, w) + D(a, w)$ becomes :

$$\frac{db_i}{dt} = \int_{\Omega} \phi_i \cdot \left(I + \sum_{p=0}^n L(\phi_p) b_p + \sum_{p,q=0}^n \left(C(\phi_p, \phi_q) + \frac{D(\phi_p, z_q)}{p} \right) b_p b_q \right)$$

ESTIMATION OF *a*

$$\forall x \in \mathbb{R}^d, \sum_{i=n+1}^N b_i(t)\phi_i(x)\Delta t = (u(x,t) - w(x,t))\Delta t$$

realization of $d\tilde{X}_t^x \stackrel{\Delta}{=} \sigma(x,t)dB_t$ (reminder : $a = \sigma\sigma^t$)

$$z_{i}(x) = \int_{0}^{T} \frac{b_{i}(t)}{\lambda_{i}T} a(x,t) dt$$

$$= \int_{0}^{T} \frac{b_{i}(t)}{\lambda_{i}T} d < \tilde{X}^{x}, \left(\tilde{X}^{x}\right)^{t} >_{t}$$

$$= \mathbb{P} - \lim_{\Delta t \to 0} \sum_{t_{i}=0}^{T} \frac{b_{i}(t_{i})}{\lambda_{i}T} \left(\tilde{X}^{x}_{t_{i+1}} - \tilde{X}^{x}_{t_{i}}\right) \left(\tilde{X}^{x}_{t_{i+1}} - \tilde{X}^{x}_{t_{i}}\right)^{t}$$

 Estimator with good statistical properties (Rao (1999)) (when the projection subset is well chosen)

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INTRODUCTION

STOCHASTIC FLUID DYNAMIC MODEL

REDUCED MODEL

RESULTS

Data Turbulence modes z_i Reconstruction of temporal modes

CONCLUSION



Data

Wake behind a cylinder

▶ at Reynolds 300 : 2D flow



INTRODUCTION	STOCHASTIC FLUID DYNAMIC MODEL	REDUCED MODEL	RESULTS	CONCLUSION	Références
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Data

Wake behind a cylinder

- ▶ at Reynolds 300 : 2D flow
- ▶ at Reynolds 3900 : 3D flow



Vorticity

TURBULENCE MODES z_i



FIGURE : Local spectral representations of the matrix $z_i(x, y, 0)$, normalized by $\sqrt{\lambda_i} = \sqrt{\overline{b_i^2}}$, from the 2D flow (Re=300).

TURBULENCE MODES z_i



FIGURE : Local spectral representations of the matrix $z_i(x, y, 0)$, normalized by $\sqrt{\lambda_i} = \sqrt{\overline{b_i^2}}$, for some some point (x, y, 0) of the horizontal section at z = 0, from the 3D flow (Re=3900).





FIGURE : n=2 and Re=3900 (3D flow).

Increase of a estimation :

0

$$\frac{\partial w}{\partial t} = I + L(w) + C(w, w) + D(a, w) \text{ becomes :}$$

$$\frac{db_i}{dt} = \int_{\Omega} \phi_i \cdot \left(I + \sum_{p=0}^n L(\phi_p) b_p + \sum_{p,q=0}^n \left(C(\phi_p, \phi_q) + \alpha D(\phi_p, z_q) \right) b_p b_q \right)$$





FIGURE : n=2 and Re=300 (2D flow).



FIGURE : n=2 and Re=3900 (3D flow).

Modal increase of a estimation :

$$\frac{\partial w}{\partial t} = I + L(w) + C(w, w) + D(a, w) \text{ becomes}:$$

$$\frac{db_i}{dt} = \int_{\Omega} \phi_i \cdot \left(I + \sum_{p=0}^n L(\phi_p) b_p + \sum_{p,q=0}^n \left(C(\phi_p, \phi_q) + \alpha_{iq} D(\phi_p, z_q) \right) b_p b_q \right)$$



FIGURE : n=6 and Re=3900 (3D flow).

PLAN

INTRODUCTION

STOCHASTIC FLUID DYNAMIC MODEL

REDUCED MODEL

RESULTS

CONCLUSION

SUMMARY

- Building of a complete physical fluid dynamic model with few assumptions
- Local variance-covariance matrix *a*(*x*, *t*) projected on large-scale velocity POD-modes :
 Simple autonomous reduced model and physical analysis of turbulence
- Direct application on POD-Galerkin model : Encouraging results
 But, the time-decorrelated part of the small-scale random velocity seems not enough energetic
- Corrective coefficient or matrix : Simple method with great results

FURTHER WORK

- w semi-martingale : which implies non-Gaussian and partially temporally-correlated random small-scale velocity
- Stochastic oceanic models : Geostrophic, Shallow water, Boussinesq, QG, SQG, ...
- Estimation of *a* by Kriging estimation of measurements errors (SSH)

QUESTIONS

Thank you for your attention

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